

# SEASONAL VARIATION OF THE LUNAR TIDAL EFFECTS IN THE $F_2$ LAYER OF THE IONOSPHERE OVER INDIAN STATIONS

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## ABSTRACT

The lunar semi-diurnal oscillations in the midday values of the critical frequency ( $f_0F_2$ ) and the height of maximum electron density ( $h_pF_2$ ) of the  $F_2$  layer are computed for all Indian ionospheric stations separately for each season of the year. The amplitude of oscillation in  $f_0F_2$  is found to be larger in winter than in summer at each of the stations. There is a reversal in the phase of the oscillation in  $f_0F_2$  between the equatorial and tropical latitudes and this is most evident in the winter months and is almost absent in summer. The annual average oscillation in  $f_0F_2$  is in agreement with that found in a previous paper (Rastogi, 1961). The phase has a large seasonal variation of about  $180^\circ$  at an equatorial or a tropical latitude station. The phase and amplitude of the lunar tide in  $h_pF_2$  do not vary significantly with latitude or with season.

## INTRODUCTION

THE existence of large lunar tidal effects in the  $F_2$  layer of the ionosphere was established by the analyses of data at Huancayo (Martyn, 1947; McNish and Gautier, 1949 *a*), Canberra (Martyn, 1948), Slough (Appleton and Beynon, 1948), and at a few other places (Martyn, 1949 *b*).

Geomagnetic control of the phase of the lunar semidiurnal oscillation  $L_2$  in the noon critical frequency  $f_0F_2$ , was first demonstrated by McNish and Gautier (1949 *b*). At stations close to the geomagnetic equator, the noon  $f_0F_2$  was shown to be maximum about two days after the first and last quarters of the moon, *i.e.*, at 04 and 16 lunar hours, while at about  $20^\circ$  geomagnetic latitude, the maxima were shifted about  $180^\circ$  in phase. These conclusions have been confirmed by later analyses (Osborne, 1952; Brown, 1956; Kotadia and Ramanathan, 1956).

Rastogi (1961 *a, b*) showed by an analysis of the lunar tide in  $f_0F_2$  at a large number of stations that the phase of the tide in  $f_0F_2$  at any station is determined by its magnetic rather than by its geomagnetic or geographic latitude. The reversal of phase from the equatorial type of variation with a maximum near 04 lunar hour to the higher latitude type with a maximum near 10 lunar hour occurs at about  $\pm 11^\circ$  magnetic latitude. The latitudinal variation of the amplitude shows a sharp maximum on the magnetic equator and two broad maxima at about  $\pm 20^\circ$  magnetic latitude.

The lunar perturbations in  $f_0F_2$  at low latitudes are closely associated with the development of the anomalous equatorial belt of the  $F_2$  layer (Rastogi, 1963). As there has been a chain of ionospheric stations in the Indian zone covering magnetic latitudes  $02^\circ$  to  $25^\circ$ , it seemed useful to study the lunar tide in  $F_2$  at these stations in greater detail. In this article, the seasonal variation of the lunar tide in the midday values of the  $F_2$  parameters are studied as a first step in this direction.

#### MATERIAL FOR ANALYSIS

The data utilised in this paper are mainly the midday (11–13 hour mean) values of the critical frequency,  $f_0F_2$ , and the height of maximum electron density  $h_pF_2$  as determined by the virtual height at a frequency equal to 0.834 times  $f_0F_2$ , for all Indian ionospheric stations during the period of low solar activity. The data used here cover the period January 1951 to December 1955 for all the stations other than Ahmedabad for which the data from March 1953 to February 1956 were used. The method of computation of the amplitude, phase and their probable errors are described in earlier papers (Rastogi, 1961 *b*, 1962).

#### SEASONAL VARIATION OF LUNAR TIDE IN MIDDAY VALUE OF $f_0F_2$

The variation with lunar phase of the deviation of midday value of  $f_0F_2$  from its monthly mean value averaged for winter (Nov., Dec., Jan. and Feb.), equinoxes (March, April, Sept. and Oct.), summer (May, June, July and Aug.) and the whole year are shown in Fig. 1 for each of the stations. The thick continuous line shows the 12 lunar-hourly wave derived from the Fourier analysis of the average deviation of  $f_0F_2$  against lunar phase. The average values of  $f_0F_2$  as well as the amplitudes (with probable errors) and phases of lunar semidiurnal variation in  $f_0F_2$  for different seasons are given for each station in Table I. The coefficients of lunar semidiurnal variation at different stations are plotted in harmonic dial in Fig. 2 separately for each season,

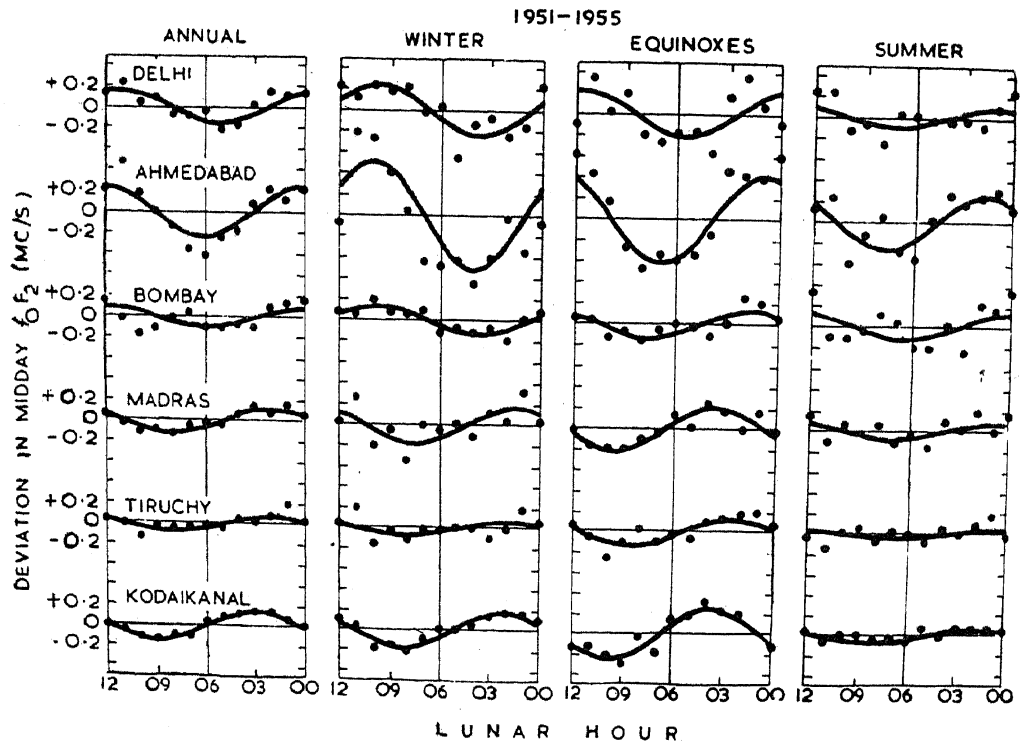


FIG. 1. Lunar tidal variations in the midday value of  $f_0F_2$  at Indian stations during different seasons and average for the whole year.

Referring to the annual curves in Fig. 1, the shift of phase of lunar variation between the equatorial stations Kodaikanal, Tiruchy and Madras and the higher latitude stations Bombay, Ahmedabad and Delhi is clearly seen. Within the equatorial group of stations, the largest amplitude occurs at Kodaikanal which is nearest the magnetic equator. Similarly among the higher latitude stations, the largest amplitude occurs at Ahmedabad which is the station closest to the region of maximum afternoon  $f_0F_2$ . This confirms earlier conclusions (Rastogi, 1961).

The features described above are most clearly seen in the winter months. In the equinoctial months, however, the phase change takes place gradually between the equator and higher latitudes, the time of maximum being 3.7 lunar hour at Kodaikanal and 0.0 lunar hour at Delhi.

Rather unexpectedly, it is found that during the summer months there is an absence of phase reversal between the equatorial and temperate latitudes. Referring to Fig. 2, the points referring to different stations all lie in the first quadrant of the harmonic dial, the phases varying only with in

TABLE I

Lunar semi-diurnal oscillation in midday (11-13 hr. mean) value of  $f_0F_2$  at Indian Stations during 1951-1955

Station	Annual average			Equinoxes		
	Average $f_0F_2$	Amplitude of $L_2$	Phase of max. $L_2$	Average $f_0F_2$	Amplitude of $L_2$	Phase of max. $L_2$
	Mc./s.	Mc./s.	Lunar Hr.	Mc./s.	Mc./s.	Lunar Hr.
hi ..	9.03	$0.15 \pm 0.022$	11.1	10.00	$0.22 \pm 0.039$	0.0
medabad	9.91	$0.34 \pm 0.070$	0.2	11.23	$0.50 \pm 0.070$	0.9
nbay ..	10.52	$0.09 \pm 0.019$	11.8	11.20	$0.13 \pm 0.031$	1.6
dras ..	8.45	$0.01 \pm 0.018$	2.0	8.91	$0.21 \pm 0.033$	3.6
ichy ..	8.13	$0.07 \pm 0.016$	2.0	8.45	$0.13 \pm 0.028$	2.5
laikanal ..	8.13	$0.14 \pm 0.017$	3.0	8.63	$0.25 \pm 0.032$	3.7

Winter			Summer		
Average $f_0F_2$	Amplitude of $L_2$	Phase of max. $L_2$	Average $f_0F_2$	Amplitude of $L_2$	Phase of max. $L_2$
Mc./s.	Mc./s.	Lunar Hr.	Mc./s.	Mc./s.	Lunar Hr.
8.50	$0.24 \pm 0.039$	9.7	8.60	$(0.09 \pm 0.035)$	1.0
10.00	$0.51 \pm 0.080$	10.0	8.50	$0.26 \pm 0.060$	1.7
10.70	$0.13 \pm 0.035$	9.7	9.67	$0.12 \pm 0.032$	0.4
8.60	$0.17 \pm 0.032$	1.4	7.84	$(0.07 \pm 0.028)$	0.9
8.31	$(0.06 \pm 0.027)$	2.0	7.61	$(0.04 \pm 0.027)$	0.1
8.24	$0.17 \pm 0.031$	2.2	7.51	$(0.03 \pm 0.027)$	2.3

N.B.—The bracketed figures indicate amplitudes which are not statistically significant according to Mann, 1951.

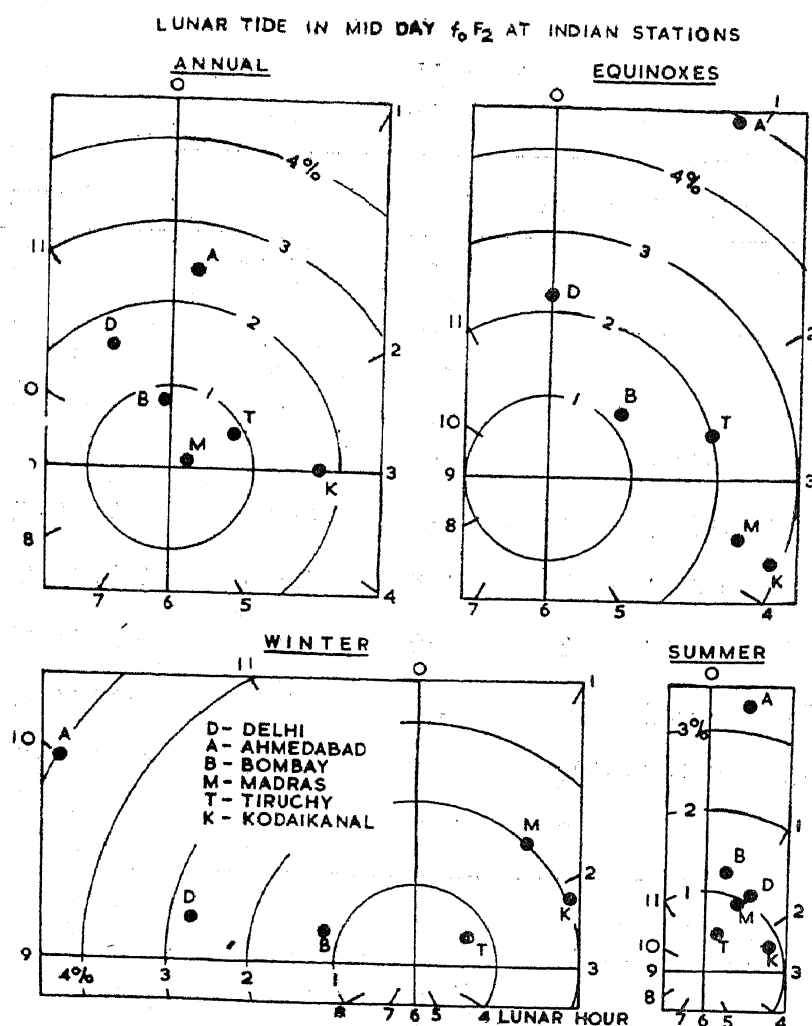


FIG. 2. Harmonic dials showing the coefficients of lunar tidal variation in midday value of  $f_0F_2$  at Indian stations during different seasons of the year.

00 and 02 lunar hour. Except at Ahmedabad, the amplitudes at other stations are small and not significantly different from each other.

Chapman (1951) has expressed the view that any determination of amplitude should be considered as statistically significant only if it is at least three times its probable error. Values of amplitude in Table I which do not fulfil this condition are shown bracketed. It is seen from Table I that most of the amplitudes are statistically significant except those for equatorial stations in summer, when the amplitude itself is very small.

The phase at temperate latitudes is about 10 lunar hour during winter and about 01 lunar hour during equinoxes and summer. At equatorial

stations, the phase is about 02 lunar hour during the winter or the summer and about 03 lunar hour during equinoxes. Thus the phase difference between the lunar variation of  $f_0F_2$  at the two groups of stations is largest (about  $120^\circ$ ) during the winter and the least (about  $30^\circ$ ) during the summer months.

To check whether the non-reversal of the phase of lunar tide between equatorial and higher latitudes summer was peculiar to Indian stations an analysis was made of  $f_0F_2$  data at the equatorial station Nhatrang and the temperate latitude station Okinawa which lie in the same zone but further east.

The lunar variations in midday value  $f_0F_2$  at Okinawa and Nhatrang in winter and summer are shown in Fig. 3. The variations are seen to be almost opposite in phase in the winter months and about 3 hours out of phase in the summer months.

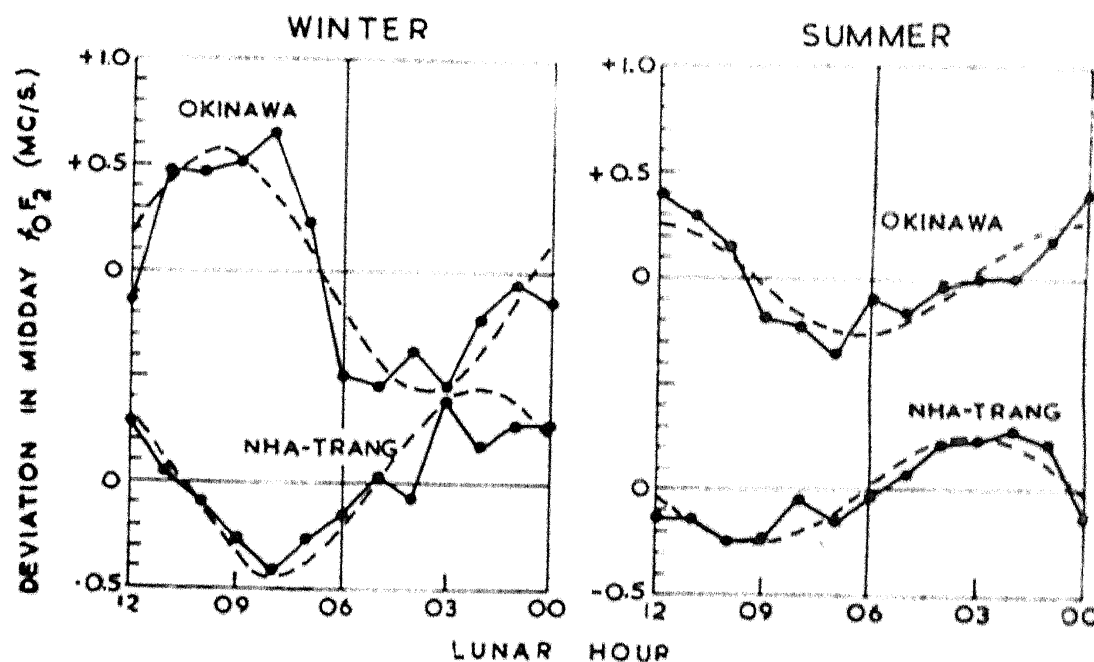


FIG. 3. Lunar tidal variations in midday value of  $f_0F_2$  at Okinawa and Nhatrang during winter and summer months.

#### LUNAR TIDE IN $f_0F_2$ AT EQUATORIAL AND TEMPERATE LATITUDE STATIONS DURING INDIVIDUAL MONTHS

To elucidate further the seasonal variation of the lunar tide in  $f_0F_2$ , an analysis was made of  $f_0F_2$  data at Ahmedabad and Kodaikanal separately for individual months. To keep up the volume of data analysed, all available

data from July 1952 to December 1959 for Kodaikanal and from March 1953 to December 1962 for Ahmedabad were utilised. The results of the analysis are plotted as harmonic dials in Figs. 4 and 5 for Ahmedabad and Kodaikanal respectively.

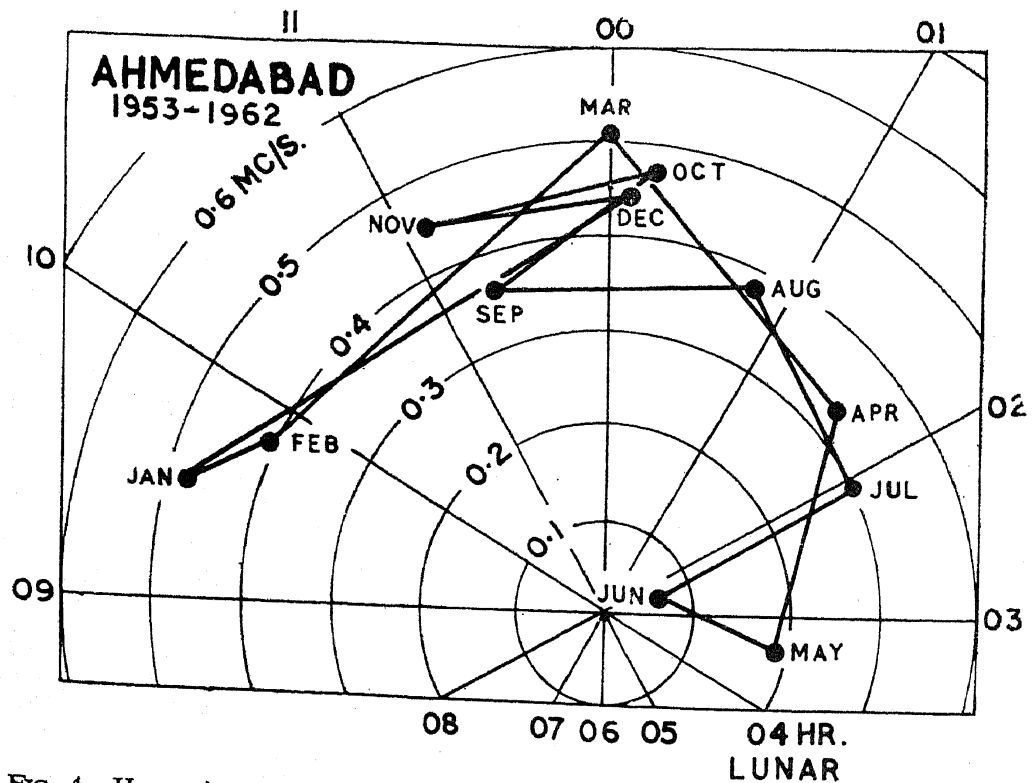


FIG. 4. Harmonic dials showing the coefficients of lunar tidal variations in midday value of  $f_0F_2$  at Ahmedabad during each month of the year.

Referring to Fig. 4 one finds that the phase of the  $L_2$  vector moves systematically through almost  $180^\circ$  during the year; being at about 9.5 lunar hour in January-February and at about 03 lunar hour in May-June. The length of the vector is greater in January than in June. The maximum amplitude occurs during equinoxes when the average value of  $f_0F_2$  itself is also the maximum.

At Kodaikanal also the phase changes almost through  $180^\circ$  varying from 0.5 and 6.5 lunar hour. The amplitude is extremely small during August and its phase is apparently at 9.0 lunar hour. These are not at all significant.

#### LUNAR TIDE IN HEIGHT OF MAXIMUM ELECTRON DENSITY IN $F_2$ LAYER

The published data of all the Indian stations except Kodaikanal include the virtual height of reflection at a frequency equal to 0.834 times  $f_0F_2$

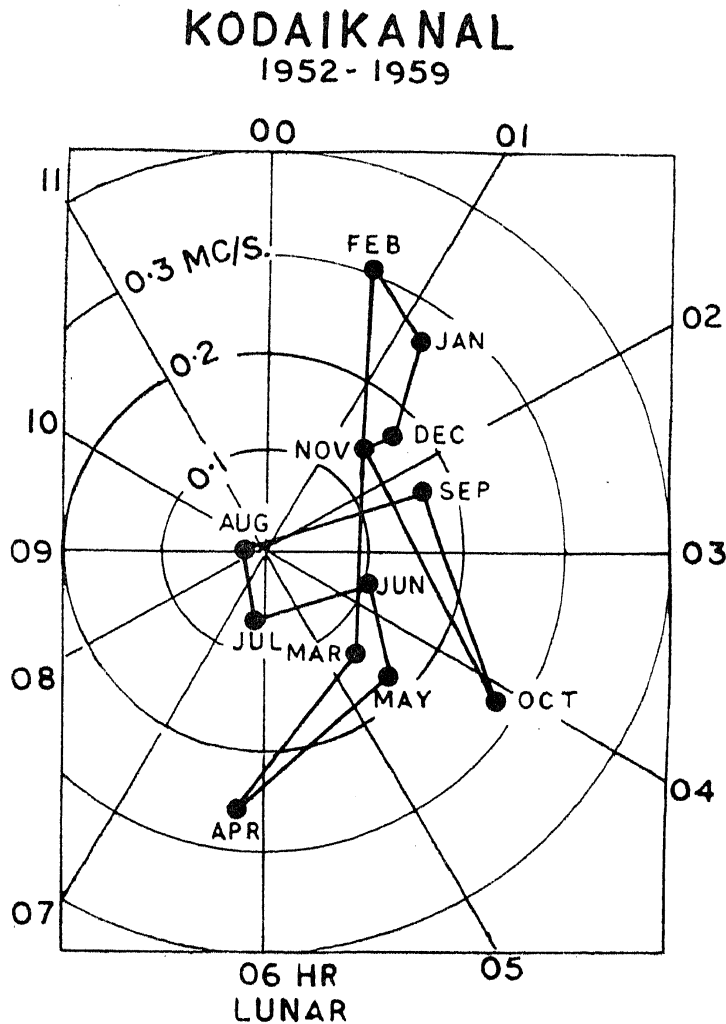


FIG. 5. Harmonic dials showing the coefficients of lunar tidal variations in midday value of  $f_0F_2$  at Kodaikanal during each month of the year.

denoted as  $h_pF_2$ . This represents the height of the maximum electron density if the shape of the layer is assumed to be parabolic. An analysis was made of the midday values of  $h_pF_2$  at each station separately for different seasons. The amplitudes and phases are given in Table II. Here again the statistically non-significant values are enclosed within brackets. The annual average deviations of  $h_pF_2$  with the lunar phase for each station are shown in Fig. 6.

Referring to Table II, one finds that the determination of amplitude at Bombay for any of the seasons or for the whole year is not statistically significant. The annual average value of the amplitude at all other stations, except Bombay, are seen to fulfil the criterion of statistical significance. The large random errors in the points for Bombay are clearly seen in Fig. 6. The



TABLE II

*Lunar semi-diurnal oscillation in midday (11-13 hr. mean) value of  $h_p F_2$  at Indian Stations during 1951-55*

Station	Annual average			Equinoxes		
	Average $h_p F_2$	Amplitude of $L_2$	Phase of max. $L_2$	Average $h_p F_2$	Amplitude of $L_2$	Phase of max. $L_2$
	Km.	Km.	Lunar Hr.	Km.	Km.	Lunar Hr.
Delhi ..	290	$2.6 \pm 0.45$	6.6	289	$2.5 \pm 0.8$	5.9
Ahmedabad	339	$4.3 \pm 0.71$	8.3	335	$4.3 \pm 1.2$	8.2
Bombay ..	405	$(0.6 \pm 0.46)$	2.5	409	$(1.5 \pm 0.8)$	5.2
Madras ..	441	$2.8 \pm 0.54$	5.9	444	$(2.1 \pm 0.9)$	8.3
Tiruchy ..	512	$2.7 \pm 0.49$	4.5	509	$(0.8 \pm 0.8)$	4.4

Winter			Summer		
Average $h_p F_2$	Amplitude of $L_2$	Phase of max. $L_2$	Average $h_p F_2$	Amplitude of $L_2$	Phase of max. $L_2$
Km.	Km.	Lunar Hr.	Km.	Km.	Lunar Hr.
260	$2.4 \pm 0.6$	6.5	322	$(2.6 \pm 0.9)$	6.8
300	$4.1 \pm 0.9$	7.2	382	$5.4 \pm 1.5$	9.1
379	$(1.8 \pm 0.7)$	10.7	428	$(1.9 \pm 0.9)$	3.6
425	$3.7 \pm 1.0$	5.5	454	$4.1 \pm 0.9$	5.4
508	$4.2 \pm 0.8$	5.3	518	$3.8 \pm 0.9$	3.6

*N.B.*—The bracketed figures indicate amplitudes which are not statistically significant according to Chapman, 1951.

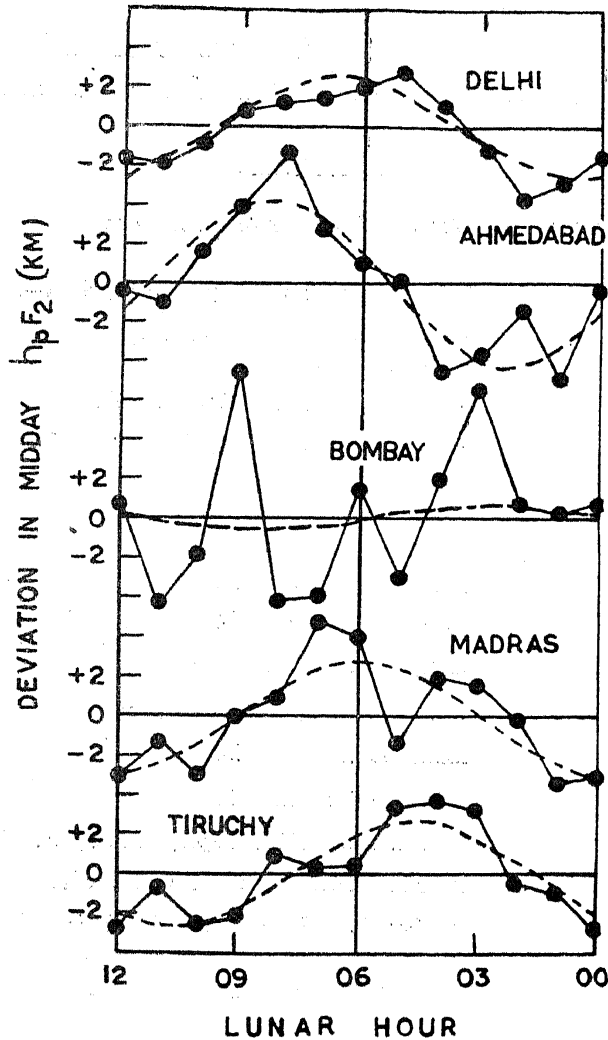


FIG. 6. Annual average lunar tidal variations in the midday value of  $h_p F_2$  at Indian stations.

lack of significance at other stations for some reasons is not unexpected because the individual data of  $h_p F_2$  are greatly affected by the retardations in the propagation of radio waves in the lower layers. The relative closeness of  $f_0 F_1$  to  $f_0 F_2$  would produce large errors in the determination of  $h_p F_2$ .

The amplitude of lunar tide in midday value of  $h_p F_2$  is about 4 Km. at Ahmedabad and about 3 Km. at other stations. The phases are not significantly different at these places being within about 05 and 08 lunar hours. This confirms the conclusions by earlier authors (Duncan, 1956) that the lunar tide in the virtual height of the  $F_2$  layer ( $h'F_2$ ) does not show any significant variation with latitude.

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## REFERENCES

- Appleton, E. V. and Beynon,  
W. J. G. *Nature, Lond.*, 1948, **162**, 486.
- Brown, R. A. .. *J. Atmosph. Terr. Phys.*, 1956, **9**, 144.
- Chapman, S. .. *Compendium of Meteorology*, American Meteorological Society, 1951, p. 510.
- Duncan, R. A. .. *Austral. J. Phys.*, 1956, **9**, 112.
- Kotadia, K. M. and  
Ramanathan, K. R. *Proc. Ind. Acad. Sci.*, 1956, **43**, 394.
- McNish, A. G. and Gautier,  
T. N. *J. Geophys. Res.*, 1949 *a*, **54**, 181.
- \_\_\_\_\_ .. *Ibid.*, 1949 *b*, **54**, 303.
- Martyn, D. F. .. *Proc. Roy. Soc.*, 1947, **190 A**, 273.
- \_\_\_\_\_ .. *Ibid.*, 1948, **194 A**, 429.
- \_\_\_\_\_ .. *Nature, Lond.*, 1949, **163**, 34.
- Osborne, B. W. .. *Ibid.*, 1952, **169**, 661.
- Rastogi, R. G. .. *Ibid.*, 1961 *a*, **189**, 214.
- \_\_\_\_\_ .. *J. Atmosph. Terr. Phys.*, 1961 *b*, **22**, 290.
- \_\_\_\_\_ .. *J. Res. (N.B.S.)*, 1962, **66 D**, 601.
- \_\_\_\_\_ .. *J. Geophys. Res.*, 1963, **68**, 1166.